

# **A combined analysis of geomagnetic data and cosmic ray secondaries for the September 2017 space weather event studies**

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**Abstract.** The September 2017 solar flares and the subsequent geomagnetic storms driven by the coronal mass ejections were recognized as the ones of the most powerful space weather events during the current solar cycle. The occurrence of the most powerful solar flares and magnetic storms during the declining phase of a solar cycle is a common phenomenon, and the current cycle is no exception. Nowadays, thorough and multifactor space weather monitoring is required to prevent the damages from the destructive space weather impact on the technological systems on the Earth. The purpose of this study is to better characterize these events by applying the generalized characteristic function approach for combined analysis of geomagnetic activity indices, total electron content data and secondary cosmic ray data from the muon hodoscope that contained Forbush decreases resulting from solar plasma

impacts. A combined analysis of secondary cosmic ray data from the muon hodoscope, geomagnetic activity indices and total electron content data is presented. The main advantage of this approach is the possibility to identify low-amplitude specific features in datasets characterizing several environmental sources. As an example, different datasets available over the storm period 6–11 September 2017 were analyzed in a unified way. The new developed technique allows us to study various space weather effects and obtain new mutually supportive information on different phases of storm evolution, based on the geomagnetic and other environmental observations in the near-Earth space.

## **1. Introduction**

The geomagnetic field protects the Earth's atmosphere from the solar wind, which consists of charged particles released by the Sun's magnetic field. The interaction between the Sun's and the Earth's magnetic fields is complex and important to understand the changes in the solar-terrestrial environment on time scales ranging from minutes to glacial cycles. Many solar obser-

variations and parameters exist, ranging from very long timescales to short timescales, however, sunspots [e.g. *Solanki*, 2003] remain the most prominent indicator of both solar magnetism and solar activity. Now the sunspot cycle has been reconstructed over millennia [*Usoskin*, 2017] and it represents a manifestation of internal processes best observed in solar magnetic phenomena. The sunspot numbers are generally correlated with both solar flares and Coronal Mass Ejections (CMEs) rate. These two classes of the largest solar eruptive phenomena are the primary causes of space weather events, we are interested here. They are frequently, if not always, linked to each other, and are the product of instabilities in the sun multi-scale dynamic magnetic structures. The dynamic processes in the Sun and its atmosphere propagate into interplanetary space to the Earth's near-space. The medium of propagation is the solar wind, a plasma that becomes supersonic as it expands from the solar corona to fill the heliosphere, as first proposed in [*Parker*, 1958], and then extensively described and documented in, e.g., [*Balogh et al.*, 2014].

Radiation and particles emitted by the Sun, with variable delays, interact with the Earth's magnetic field and the atmosphere, and cause electrical currents to

flow in regions of the ionosphere and magnetosphere. The solar wind, travelling into the space, complicates more the near-Earth environment, and mainly its magnetic environment. The Earth's magnetic field comprises contributions from sources inside the Earth (internal contributions, including those from the liquid core and lithosphere) and outside it (external contributions, including those from ionosphere, magnetosphere and their coupling). The external sources induce secondary fields on the Earth. Important to note is that the core magnetic field acts like a shield to the solar wind that the Sun continually emits. Thereafter, to understand the Sun-Earth environment evolution, the variations of the geomagnetic field need also to be understood, e.g. [*Mandea and Purucker, 2018*].

The temporal variations of the geomagnetic field are mainly described from the ground-based magnetometer measurements. Here, we are interested in rapid variations of the field, linked to the Sun-Earth interactions. They are described by the so-called geomagnetic activity indices. Among them, two are widely used to characterize the geomagnetic conditions,  $K_p$  and  $Dst$  indices. The  $K_p$ , a geomagnetic three-hour-range index, is derived from the standardized  $K$  index of 13 magnetic observatories and is intended to measure the solar

particle radiation by its magnetic effects. The *Dst* index is derived from data provided by 4 near-equatorial geomagnetic observatories and is designed to measure the intensity of the globally symmetrical equatorial electrojet (the “ring current”). *Dst* is then a measure of geomagnetic activity and is used to assess the severity of magnetic storms. For more details, the reader can access the International Service of Geomagnetic Indices (<http://isgi.unistra.fr/>).

Space weather is defined as the collection of physical processes, beginning at the Sun, and ultimately affecting human activities on Earth and in space (Natural Resources Canada, Space Weather, <http://spaceweather.gc.ca/sbg-en.php>. Accessed 07.09.2017). The geometry of the geomagnetic and interplanetary magnetic fields, and their evolution in space and time contribute to the complexities of space weather observed at the Earth’s surface, and in the near-Earth space. The electrical currents, and the changing energetic particle populations result in geomagnetic variations, aurora, and can profoundly affect a number of technologies. There are well-known examples of disastrous consequences of some of these disturbances. Induced currents caused by magnetic storms can lead to saturation, overheating and dam-

age of the high-voltage transformers in electrical substations. In this study we precisely investigate the September 2017's geomagnetic storm linked to an important solar activity observed as Martian aurora [*Xu et al.*, 2018], and across the globe on Earth. The September 2017 magnetic storm driven by the CME also was followed by multiple radio blackouts on the daylight side of the Earth (NOAA Space Weather Scales // NOAA / NWS Space Weather Prediction Center. URL: <https://www.swpc.noaa.gov/noaa-scales-explanation>. Last visited: 12.03.2018); a similar accident occurred after the storm driven by the X9.0-class solar flare in 2006. Nowadays, thorough and multifactor space weather monitoring is required to prevent the damages from the destructive space weather impact on the technological systems on the Earth [*Gvishiani et al.*, 2016a, 2016b]. To take into consideration these geomagnetic storm aspects, the remainder of the paper is organized as follows. Section 2 presents the different datasets used in the current study. In Section 3 some information on the applied method and its advantages is given, mainly to show the possibilities to apply it for rapid variations as geomagnetic storms are. Section 4 discusses a specific case study, focused on the September 2017 magnetic storm, and finally Section 5 draws

some conclusions and maps out future directions.

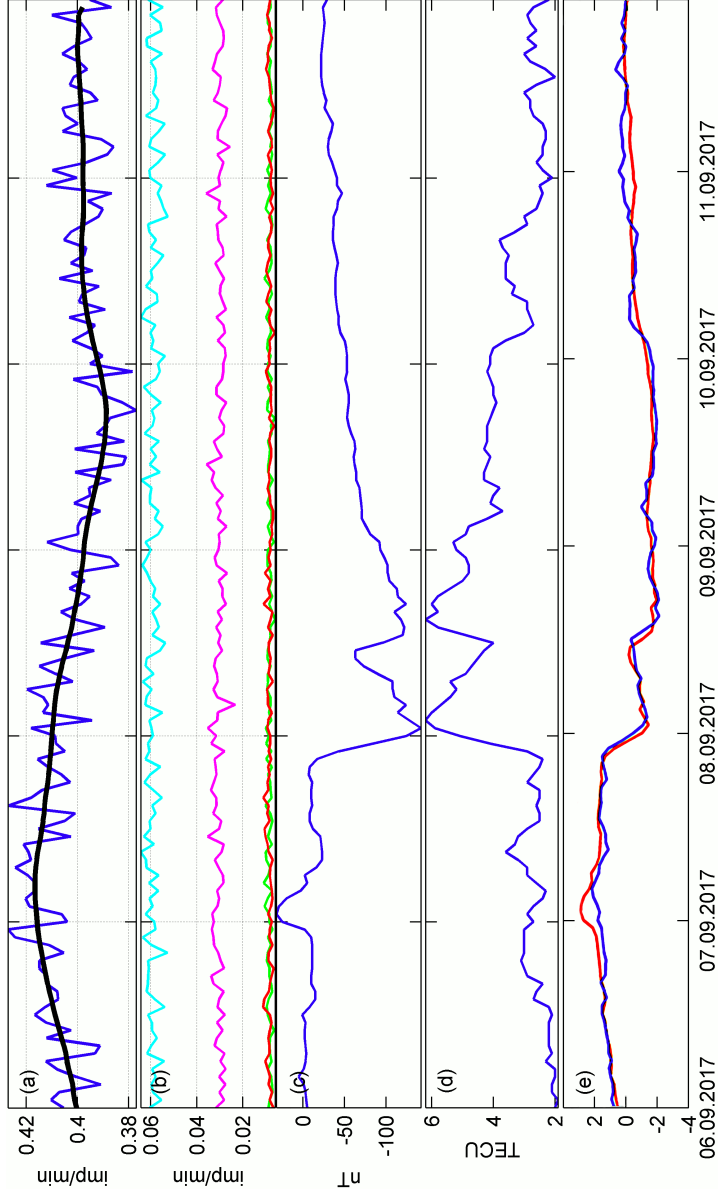
## 2. Data

In this study, we analyzed the space weather parameters and geomagnetic activity indices along with the time series from the URAGAN (“Ustanovka RAspoznavainya Grozovykh ANomaliy” – “Thunderstorm Anomaly Recognition Set” in Russian) muon hodoscope data [Barbashina et al., 2008]. Muon diagnostics of the near-Earth space is a modern technology applied in various problems of cosmic ray physics, ionospheric and geomagnetic studies. URAGAN is a wide-aperture, large-area multilayered muon hodoscope [Barbashina et al., 2008], which was designed for studying various phenomena in the circumterrestrial space and, in particular, in the Earth’s atmosphere [Barbashina et al., 2017; Mikhaylenko et al., 2011] and magnetosphere [Astapov et al., 2017], that cause variations in the muon flux at the ground level. The hodoscope provides recording of the muon flux intensity from different azimuthal directions ( $0^{\circ}$ – $360^{\circ}$ ) and zenith angles ( $0^{\circ}$ – $80^{\circ}$ ) with a high degree of spatial and angular accuracies (1 cm and  $1^{\circ}$ , respectively) [Barbashina et al., 2008]. It enables registration of each muon along with its track



reconstruction, as well as every 1-minute construction of two-dimensional angular matrix displaying the muon flux from the observed hemisphere. For the current study, we reassembled the data matrices from 1-minute to 1-hour by a simple averaging. Variations of the muon flux in the near-Earth space, which are the secondary cosmic ray data, are associated with the physical processes taking place at the time of anomalous space weather events. During the quiet periods, the muon flux intensity at a ground level is stable and relatively high. Before major geomagnetic storms driven by CME, a strong and rapid decrease can be seen in the muon intensity time series. The effect, called the Forbush decrease [*Cane*, 2000], occurs due to the screening of the cosmic ray flux by the plasma cloud heading towards the Earth's magnetosphere. The muon data from URAGAN hodoscope has been successfully used for studying the CMEs [*Astapov et al.*, 2017] and corresponding Forbush decreases.

Figure 1a depicts the muon intensity values extracted from a central cell of each data matrix from 6 September 2017, 00:00 UT to 11 September 2017, 23:00 UT. Figure 1b displays the muon intensity time series for corner cells of the matrices. The maximal intensity comes from the zenithal direction to the central area



**Figure 1.** Inter-comparison between the muon intensity data from the center of the moon hodoscope matrices (a), from the four corners of the hodoscope matrices (b), the *Dst* index (c), total electron content (d), and the generalized characteristic functions (e) G1 (red) and G2 (blue).

of the hodoscope, thus, the muon intensity data from the central cell is of the most interest.

We also used hourly sampled total electron content (TEC) data which is available at the IZMIRAN ionosphere weather database portal (<http://www.izmiran.ru/ionosphere/weather/storm/index.shtml>). In the GNSS signal processing practice, TEC represents the total number of electrons between transmitter and receiver. TEC data is widely used to characterize conditions of the ionosphere and to estimate the ionosphere impact on a radio wave signal [Garner *et al.*, 2008]. The more electrons in the path of the signal, the more it is affected. As TEC responds to varying solar electromagnetic and geomagnetic activity [De Haro *et al.*, 2002; Zolotukhina *et al.*, 2017], it can be definitively used in a comprehensive analysis of the space weather phenomena.

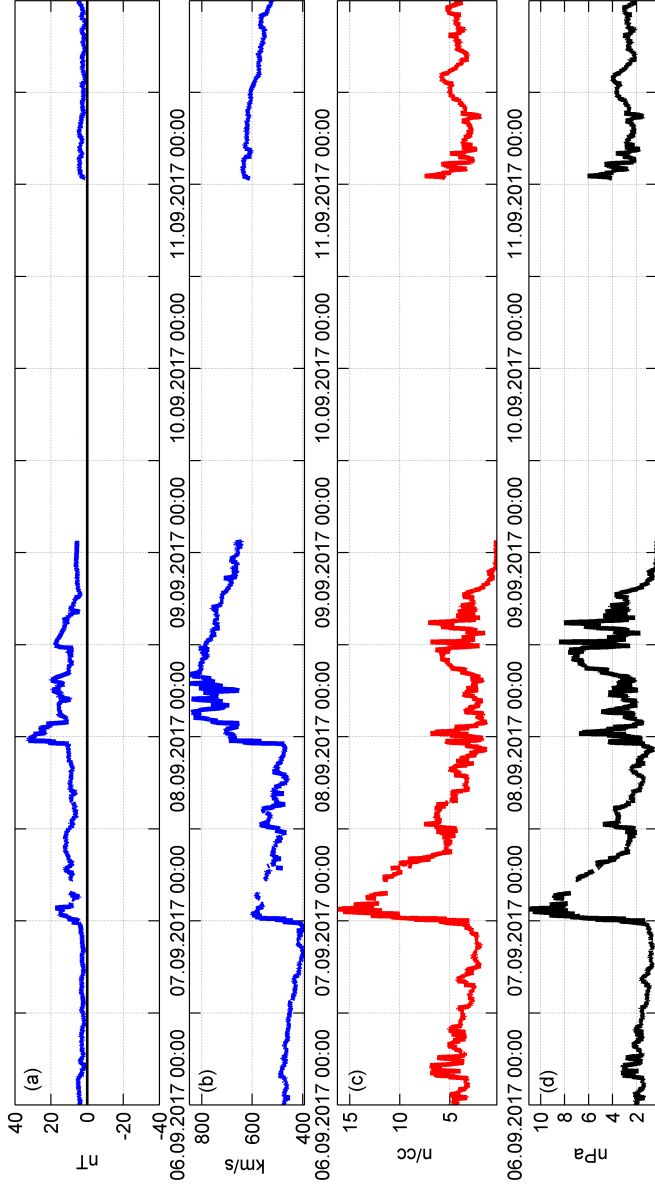
In addition, we studied *Dst* index time series over the considered period. As seen from Figure 1c and Figure 1d, the TEC time series appears to be almost a negative reflection of the *Dst* index, which indicates their similar response to the equatorial ring current.

Figure 1e displays the general characteristic functions constructed using the mentioned space weather and cosmic ray data sets. Here we place their plots so

that they can be clearly matched with the initial data. These functions will be discussed in the next sections.

Also for the demonstration of the onset and evolution of the geomagnetic storm, we used common space weather data from the OMNIWeb database (<http://omniweb.gsfc.nasa.gov>), such as the interplanetary magnetic field (IMF), the solar wind speed, density and dynamic pressure for the case study period of 6–11 September 2017. The IMF and solar wind data are 5-min-averaged. These data are plotted in Figure 2. Unfortunately, the data are missing for 9–10 September and also there are small gaps of unknown origin in the data time series at the beginning of 7 September. The gap in the solar wind data from 03:40 to 05:20 UT is most likely caused by some failure of the particle analyzer.

As the maximal  $Kp$  index values during the storm were 8+ according to the GFZ Potsdam data (<http://www.gfz-potsdam.de/sektion/erdmagnetfeld/daten-produkte-dienste/kp-index/>), the storm was classified as “severe” (G4 rate) according to the NOAA space weather scale (<http://www.swpc.noaa.gov/noaa-scales-explanation>). The first local extremum of  $Kp$  reached 8– at the end of 7 September; after that  $Kp$



**Figure 2.** Solar wind parameters during the solar flares and geomagnetic storm 6–11 September 2017: interplanetary magnetic field (a), solar wind speed (b), proton density (c) and plasma pressure (d).

decreased to its minimal value 0o on 9 September between 18:00 and 21:00. According to the 5-level classification scale based on the *Dst* index [*Loewe and Prölss*, 1997], this storm was classified as “strong”, as the minimum *Dst* level was  $-142$ .

### 3. Method

The proposed technique aims to offer a combined analysis of various space weather related data, including their morphology and time-dependent features. For the needs to better interpret the observations, it is important to identify the particular features of a physical signal, sometimes difficult to reveal due to their low amplitude. One of the techniques that can be used in this case is based on the complex indicator function (CIF) approach, previously widely used in exploration geophysics (e.g. for complex interpretation of geoelectric and geomagnetic survey data in the ore deposit search; see [*Hmelevskoy*, 1999; *Nikitin*, 1979, 1986; *Troyan and Kiselev*, 2010]), with different modifications. Here, we apply a new technique based on this approach in the space weather data analysis. As the following method, similar to CIF, however does not fully represent a classical CIF, we refer to it as a generalized characteristic

function.

The generalized characteristic function is dimensionless and thus it enables the qualitative comparison of data of different origin, dimensions and scales (ranges of variability) in a single scale of magnitudes. A common formula representing such function for some data sets  $f_1, f_2, \dots, f_n$  over the same time period or the same spatial coordinates is the following (1):

$$F(f_1, f_2, \dots, f_n) = \sum_{i=1}^n a_i f_{s_i} \quad (1)$$

where  $f_{s_i}$  is the standardized data set for an initial data set  $f_i$ , and  $a_i$  are the weight coefficients depending on the properties of a particular data set, its physical origin and reliability;  $i = 1, \dots, n$ . In our research we assumed that  $a_i = 1$ , i.e. all data sets have equal reliability.

Construction of the generalized characteristic function requires determination of typical distribution laws for all datasets under consideration. For example, some physical data can be distributed normally, and some can have the lognormal distribution law, which is defined by mean and standard deviation of the value logarithms. The standardization procedure requires the computa-

tion according to (2):

$$f_s = \frac{f - \bar{f}}{\sigma_f} \quad (2)$$

In this formula,  $\bar{f}$  is the mean  $f$  value for a whole data set, and  $\sigma_f$  is the standard deviation. In case of a lognormal distribution for  $f$ , the  $\bar{f}$  and  $\sigma_f$  values are calculated for the  $\ln f$  values.

Several criteria allow establishing the coincidence or inconsistency between estimated distribution law of selected data set and a priori known (reference) distribution law. They are calculated from the corresponding distribution function values and then estimated against a specified threshold. We used the Kolmogorov-Smirnov (KS) test [*De Smith, 2015*], which enables comparison of the two distributions using their maximum deviation. At first, the distribution of the analyzed data set  $X = (x_1, \dots, x_N)$  is built. The range of the  $X$  variability is divided into  $A$  intervals with a width  $W_x = (\max(X) - \min(X))/\sqrt{N}$  each. Then, the cumulative distribution function is calculated. For two data sets  $X = (x_1, \dots, x_N)$  and  $Y = (y_1, \dots, y_N)$ , the KS test value is determined according to the following (3):

$$D_{XY} = \max_{1 \leq i \leq A} |c_{Xi} - c_{Yi}| \quad (3)$$



where  $c_{X_i}$  and  $c_{Y_i}$  are the cumulated occurrences for the analyzed data sets  $X$  and  $Y$ , respectively. Given a data set of more than 40 values and 5% error probability, (4) represents the estimate for the critical value  $D_C$  of the KS (Estimation of a distribution law of geophysical data. Identification of background and anomalous areas using a data from a complex of methods. Shevnin V. A., Modin I. N. (eds.) The Workshop on the interpretation of electrical prospecting data, URL: [http://geophys.geol.msu.ru/STUDY/4KURS/STAT/Stat\\_.htm](http://geophys.geol.msu.ru/STUDY/4KURS/STAT/Stat_.htm), in Russian. Last visited: 12.03.2018); [*Troyan and Kiselev*, 2010]:

$$D_C = \frac{136}{\sqrt{N}} \quad (4)$$

If  $D_{XY} \leq D_C$  then distribution laws of two data sets are considered coincident. Therefore, in order to compare calculated  $D$  value and the critical value of the KS test  $D_C$ , it is needed to generate a reference cumulative distribution function using the specified data intervals, mean and standard deviation. As a result, we decided whether the distribution law of the selected data set and the reference distribution law are coincident or not.

The described approach to the analysis for multiple datasets of different origin is close to widely im-

plemented techniques based on correlation calculations between different datasets. Multiple correlations are more focused on closeness of the patterns in the analyzed data, whereas the generalized characteristic functions approach provides a data combination with a reference to normal (background) levels for each dataset. Also one of the main advantages with respect of other analysis methods is a possibility of identification of low-amplitude patterns related to the specific features of a physical process in all the analyzed datasets, using data from several environmental sources. Another significant advantage is its flexibility: it is possible to adjust the dataset contributions using weight coefficients with the standardized time series. As mentioned earlier, depending on the specialties of some additive components, they can be included into the generalized function with a negative sign, which enables setting several characteristic functions instead of one and further analysis of their mutual behaviour.

With the application of the generalized characteristic function approach to such particular space weather parameters as geomagnetic *Dst* index, TEC and cosmic ray secondary data, we highlight the particular advantages of this technique that can be useful in future space weather studies.

## 4. Case Study (September 2017 Geomagnetic Storm)

The occurrence of the most powerful solar flares and magnetic storms during the declining phase of a solar cycle (including the current 24th cycle) is a well-known phenomenon. The September 2017 solar flares and the subsequent geomagnetic storms driven by the coronal mass ejections were recognized as some of the most powerful space weather events during the current solar cycle. Such major events provide an opportunity for a study of geomagnetic activity along with the secondary cosmic ray monitoring data. The overall evolution of the 6–11 September 2017 space weather events included several steps. The first solar flare on 5 September occurred at 09:10 UTC. It blasted from a large sunspot on the Sun's surface and was rated X2.2. X-class solar flares are of the most power, according to the solar storm scale of the NOAA Space Weather Prediction Center (NOAA Space Weather Scales // NOAA / NWS Space Weather Prediction Center. URL: <https://www.swpc.noaa.gov/noaa-scales-explanation>. Last visited: 12.03.2018). The next flare at 12:02 GMT was rated even stronger – X9.3, and it was the biggest X9 flare since an X9.0 event in 2006 [*Struminsky and*

*Zimovetz*, 2010].

As clearly seen from the IMF time series (Figure 2), the storm sudden commencement occurred near 00:00 UTC on 6 to 7 September. At that time, the solar wind speed increased from 400 km/s to about 600 km/s, and there was also an increase in both proton density and flux pressure. The onset of the storm was due to the CME from the second solar flare, which reached the magnetosphere at the end of 7 September. This CME produced (1) a higher solar wind speed, which increased over 800 km/s at the beginning of 8 September (Figure 2b), and (2) a spike-shaped increase of proton density and plasma flux pressure at about 00:00–02:00 UT on 8 September (Figure 2c and Figure 2d).

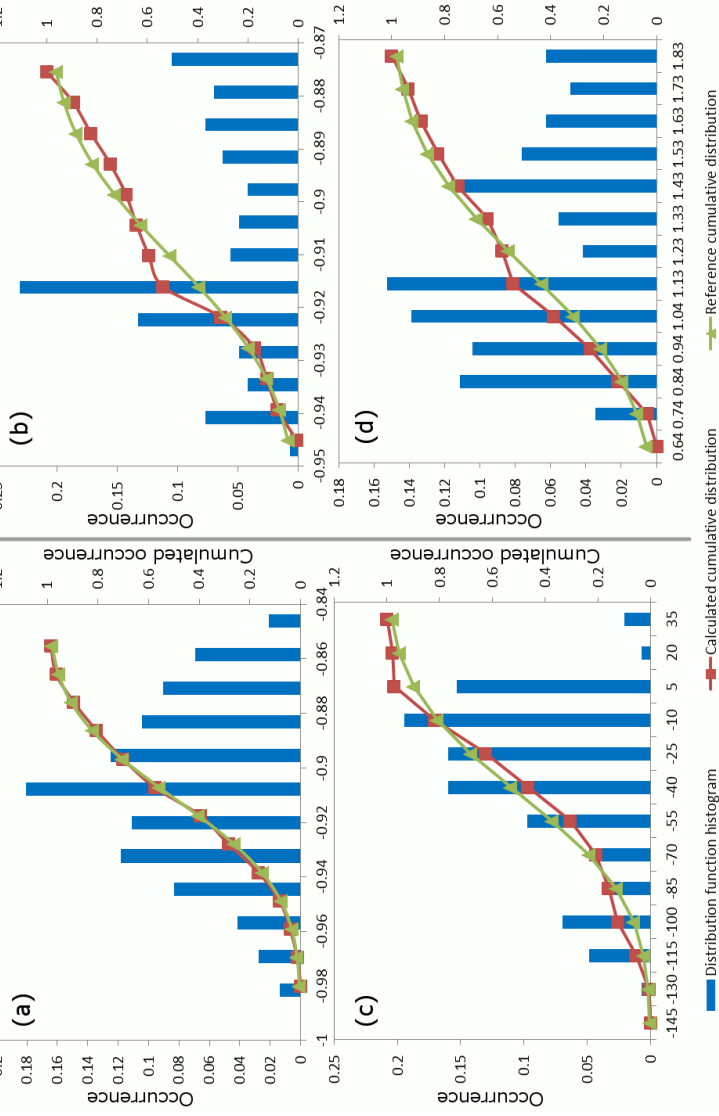
In this research, we used hourly sampled muon data corrected for the variations of pressure and temperature of the Earth's atmosphere in order to reduce the variability of muon intensity. However, it was still contaminated by variations, and therefore the trend for this time series was built using a piecewise-linear approximation. We used the approximation technique based on the local sliding approximation models [*Getmanov et al.*, 2015]. For a given time series  $y(t_i)$ , the sliding local intervals are defined:  $N_{1j} = N_d(j - 1)$ ,  $N_{2j} = N_{1j} + N - 1$ , where  $N$  is a size of the local interval, and

$N_d$  is a sliding step. On these intervals, local piecewise-linear models are defined as  $s_j(c_j, t_i) = c_{1j} + c_{2j}t_i$ . The approximation result is the sum of these models  $s_j$ , and the  $c_{1j}$ ,  $c_{2j}$  optimal coefficients for the models are determined from the quadratic functional minimization. Black curve in Figure 1a represents the resulting trend component.

Initial inter-comparison between muon,  $Dst$  index and TEC data shows that the beginning of the Forbush decrease is close to the storm sudden commencement (SSC). It is clearly seen in  $Dst$  plot as an uplift (Figure 1c) and in TEC plot as a decrease at the very beginning of 7 September. During the period of the storm onset and main phase, the Forbush decrease trend is close to linear. Very slight and insignificant oscillations of the muon flux intensity trend are seen on September 8, during two local extremums of the  $Dst$  index ( $-142$  and  $-122$  nT) and TEC (both above 6 TEC units) time series at the beginning and at the end of the day. The muon intensity trend (black curve in Figure 1a) remains quasi-constant until the 2nd half of 9 September when the recovery phase of the storm was approaching to the end. Therefore, the impact of the CME caused by the second X9.3 solar flare appeared to be not so strong comparing to the impact of the first CME.

TEC increase on 10 September, 00:00 UTC, and its consequent abrupt decrease in the morning of 10 September are practically not reflected in  $D_{st}$  values, which steadily increase to zero. At the same time, the muon intensity begins to increase 2–3 hours before 10 September midnight. It reflects the processes that can definitely have the same origin related to the change of conditions in the ionosphere due to the decay of the geomagnetic disturbance. This issue demands further research in order to study the opportunity of application of muon intensity data to adjust determination of time moments of geomagnetic storm decay.

Figure 3 shows the results of the distribution law determination for all used data sets. We checked the assumptions of normal and lognormal distributions. As the data lengths are equal, due to their equal sampling, the  $D_C$  value for the KS is 11% (equation (2)). The KS test shows that the muon flux data from the central cell of the URAGAN hodoscope has a distribution very close to lognormal (Figure 3a) with  $D = 2\%$  against  $D_C = 11\%$ . The KS test for the trend component of the muon flux intensity time series (Figure 3b) shows less coincidence (14%), however the KS test for a normal distribution of a muon trend shows an even worse coincidence (more than 20%). Therefore, we accepted



**Figure 3.** Statistical calculations to determine the distribution laws for the moon flux (a), moon flux trend component (b), *Dst* index (c) and TEC (d) datasets. The distribution function histogram (blue bar plots) and the cumulative distribution (red curve with square markers) comparing to the generated reference cumulative distribution (green curve with triangle markers) generated, are presented for each dataset. The left vertical axis shows the distribution function values, and the right vertical axis shows the cumulative distribution function values for each plot.

the lognormal distribution for the muon intensity trend. The  $Dst$  index values appear to be distributed normally with  $D = 7.3\%$  (Figure 3c), and the KS test for the TEC data set suggests its lognormal distribution, although the  $D$  value is equal to critical ( $D = 10.8\%$ ) (Figure 3d). Therefore, the generalized characteristic function should include the original values of  $Dst$  index and the natural logarithms for muon intensity and TEC data.

In this research we calculated the means and the standard deviations for the whole observation period. As mentioned before, the  $Dst$  index time series has a normal distribution. Despite the standardized values are dimensionless, they still keep the initial time series morphology and therefore can be included in the generalized function as additive components. For the time series having lognormal distribution (muon intensity and TEC data), the mean and the standard deviation are calculated from the value logarithms.

Therefore, we used two generalized functions in order to analyze the mutual relation between 1) the cosmic ray secondary data and geomagnetic data; and 2) the cosmic ray secondary data and TEC data. Mutual analysis of two generalized functions provides a better identification of possible fragments of interest, i.e. their



deviations from each other. In this study, we assumed that the muon flux data,  $Dst$  index and TEC data are of the same confidence, and, therefore, the weight coefficients for all the standardized time series are taken 1. However, it is clearly seen that the TEC time series morphology commonly repeats the  $Dst$  time series with a negative sign, which suggests taking weight coefficient -1 for standardized TEC data. The resulting generalized characteristic function is defined following (5):

$$G_1 = Mu_s + Dst_s, \quad G_2 = Mu_s - TEC_s \quad (5)$$

where  $Dst_s$ ,  $Mu_s$  and  $TEC_s$  are the standardized  $Dst$  index, standardized logarithms of muon intensity trend and TEC values, respectively.

Figure 1e shows the mutual behavior of  $G_1$  and  $G_2$  functions. The generalized characteristic functions represent the overall dynamics of muon intensity,  $Dst$  index and TEC. The behaviour of the functions is similar with the maximum Pearson correlation coefficient value for  $G_1$  and  $G_2$  is 0.9553; however, it is clearly seen that correlation can be lower in some areas. Therefore, such an approach can provide features of space weather disturbances derived from geomagnetic, ionosphere and secondary cosmic ray data. Weight coefficients for particular time series, included into the generalized char-

acteristic functions, can be customized in order to reveal specific patterns of their behaviour, such as uplifts, mistiming and others.

Commonly, the proposed approach requires no preliminary processing, like autocorrelation elimination, before the distribution laws determination. In our case, all datasets refer only to the period 6–11 September 2017. It seems that the smoothed muon flux intensity data (as well as the raw data) contain no significant diurnal variation contribution. Possible autocorrelation of  $Dst$  index values due to the contribution of trend or cyclical components is negligible, as the  $Sq$  variation is commonly eliminated during the index calculation, and the low-frequency secular variation does not affect the 5-day period data. The same refers to TEC time series, result of multiple ionosphere data processing.

## 5. Discussion and Conclusions

As seen from the data plots for the whole observation period, the muon intensity trend remains unchangeable and close to linear during the main phase and recovery phase of the geomagnetic storm. TEC data commonly represents the negative reflection of the  $Dst$  index, with

a response to the storm sudden commencement and impacts from the coronal mass ejections. Generalized function  $G_2$ , which characterizes mutual behavior of muon and TEC data, shows a series of slight uplifts closer to the end of the storm. The presence of these uplifts indicates a change in the ionosphere, due to the decay of the geomagnetic disturbance. This particularity however is not represented in the generalized function  $G_1$ , which characterizes mutual behavior of muon data and  $Dst$  index, although both functions reflect a slight muon flux intensity increase during 10–11 September.

Herein, we constructed two functions because  $Dst$  morphology repeats TEC morphology with a negative sign. Setting up two synchronized functions helped us to reveal typical patterns of the time series on the basis of mutual deviations of the generalized characteristic functions from each other.

Finally, let us summarize the main findings of this study. Firstly, we introduced the generalized characteristic functions (a modification of complex indicator function approach) for integral qualitative estimation of space weather disturbances derived from geomagnetic, ionosphere and secondary cosmic ray data. We calculated these functions over the period of 6–11 Septem-

ber 2017 in order to test our approach in the geomagnetic storm conditions. The results show that the proposed approach can be used for identification of one or more patterns representing the same physical process in space weather. Also, as mentioned, the muon flux intensity increase during the decay phase of the storm can be related to the change of conditions in the ionosphere. Fragments of incoherency between the  $G_1$  and  $G_2$  at different phases of the storm duration indicate the relative contributions from the parameters analyzed and their mutual alternation.

Summarizing all the results, the generalized characteristic functions approach is a combined analysis tool for different space weather datasets that can provide mutually supportive information on the onset, evolution and recovery phase of a geomagnetic storm. It is most likely applicable to estimating separate contributions of geomagnetic, ionosphere and cosmic ray origins during evolution of geomagnetic storm. Moreover, this approach provides an opportunity to apply muon and TEC data as a support for the determination of time moments of geomagnetic storm decay. The creation and update of storm catalogues is often performed using geomagnetic activity indices. Various circumterrestrial space monitoring tasks can require even

1-minute temporal resolution, and most of the classical indices provide a resolution of 1 hour and less. Therefore, an improvement of the storm cataloguing procedure is needed, using additional environmental data as a support. We consider that the needed efficiency can be achieved by the analysis of 1-minute sampled geomagnetic data over a set of geomagnetic observatories (or some geomagnetic index data of resolution higher than 1 hour, such as 1-min SYM/ASY indices) in combination with 1-minute muon flux intensity time series. This research describes a primary result. The method should be tested on data registered during quiet geomagnetic activity periods and on other geomagnetic storm events to improve the results and make further conclusions. Along with the development of quantitative measures of generalized characteristic functions for estimating the anomalous fragments on space weather data time series, this will be the purpose of our future studies.

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provided by the National Research Nuclear University MEPhI, and total electron content data provided by IZMIRAN ionosphere weather portal. This work employed facilities and data provided by the Shared Research Facility “Analytical Geomagnetic Data Center” of the Geophysical Center of RAS (<http://ckp.gcras.ru/>). The research has been conducted in the framework of the Russian Science Foundation project No. 17-17-01215.

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