Geomagnetic effect of the Bering Sea meteoroid

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Possibilities of studies of the geomagnetic effects produced by the interaction of a cosmic bodies with the magnetosphere-ionosphere-atmosphere system are very limited due to extremely small number of examined events. Here we present geomagnetic observations at an array of magnetometers during Bering Sea Bolide event on December 18, 2018 when a space body entered the Earth's atmosphere and exploded at the altitude of ~ 25 km near Kamchatka. It has been found that the short-lived electromagnetic signal appeared before the explosion and, consequently, was trigged by the passage of a meteoroid through the inner magnetosphere. Geomagnetic disturbances of the same duration and frequency of oscillations were detected both in the area adjacent to the explosion site in the Northern hemisphere and in the magnetically conjugated area in the Southern hemisphere. These observations may be provisionally interpreted as a triggered excitation of resonant field line oscillations in the inner magnetosphere by the fast-moving meteoroid. The magnetosphere is often in a metastable state, when even a weak external trigger can stimulate an internal instability and wave generation. The appearance of a diamagnetic effect during partial ablation of a meteoroid could cause a local disturbance of the geomagnetic field and its propagation in the magnetic force tube. KEYWORDS: meteoroid; geomagnetic response; geomagnetic pulsations; ionosphere; magnetosphere.

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Keypoints

- 1. On December 18, 2018 the Bering Sea Bolide entered the Earth's atmosphere vertically and exploded at altitude of ~25 km with a probable yield 4–35 kT.
- 2. Geomagnetic oscillatory response with frequencies 25–35 mHz was recorded \sim 3–10 min before the bolide explosion, during its passage though the inner magnetosphere.

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3. The geomagnetic response is probably due to trigger excitation of resonant field line oscillations.

Introduction

An interaction of conductive body with plasma is accompanied by excitation of various electromagnetic disturbances and emissions. In the terrestrial ionosphere these effects are manifested as generation of geomagnetic impulses and noise by take-offs of rockets [Chernogor, 2013], plasma injection in rocket experiments [Gavrilov et al., 2003], movement of large-scale space systems (e.g., tether, Shuttle orbiter, International Space Station, etc.) in the upper ionosphere [Dobrowolny and Veltri, 1986]. This class of phenomena comprises the geo-

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magnetic effects caused by the fall of large cosmic bodies, meteor showers, and bolides [Savchenko, 1975, 1976].

Studies of geomagnetic disturbances associated with the fall of meteoroids, fireballs, and meteor showers have been conducted for several decades [Bronshten, 1991]; Kalashnikov, 1949, 1952] was the first who reported the observation of weak changes of the geomagnetic field attributed to meteors. However, many subsequent studies gave no evidence of a magnetic effect due to a single meteor [Hawkins, 1958]. Observations with a high sensitivity magnetometer concluded that most individual meteors do not have any associated geomagnetic pulsation activity, but some of the larger meteors do produce magnetic effects [Ellyett and Fraser, 1963].

The greatest number of studies is devoted, perhaps, to the geomagnetic effect of the Tunguska phenomenon [Ivanov and Medvedev, 1965; Nemchinov et al., 1999. Although this event has been known for a long time, the responsible physical mechanism has not been firmly established [Bronstein, 2002; Ivanov, 1967. Geomagnetic fluctuations with quasi-period 5-13 min were observed during the Chelyabinsk meteoroid event, but ~ 45 min before the explosion [Chernogor, 2018]. Rare information on geomagnetic effects of cosmic objects hints that transients or pulsations of geomagnetic field can be generated by relatively small cosmic bodies (even with a diameter of the order of 1 m) and propagate over distances of several thousand kilometers [Chernogor, 2018]. Bodies with such dimensions invade the atmosphere quite often – once per week or month, whereas a space object commensurate with the Tunguska body falls to the Earth on average once every 100-200 years [Brown et al., 2002]. Because the impact of bolides and meteoroids on the near-Earth environment is a rare and unpredictable event, the observational information is very limited.

A large class of ionospheric and geomagnetic perturbations associated with meteor/bolides is caused by perturbation of the ionosphere by an acoustic wave coming from the region of the main energy release caused by the object destruction. These effects are physically similar to electromagnetic/ionospheric effects associated with strong ground or atmospheric explosions [Zetser et al., 2004]. Owing to waveguide effects in the atmosphere the acoustic waves from explosion can propagate to distances

about several thousand km [Adushkin et al., 2004]. Acoustic waves can modulate the plasma density and electric current in the conductive E-layer of the ionosphere, thus causing a geomagnetic response on the ground.

Besides that, other mechanisms of magnetic effect produced by bolides/meteoroids are feasible. For example, the interaction of the geomagnetic field with the plasma formed in the head of the meteoroid may be significant for geomagnetic response [Bronstein, 2002]. However, in-depth studies of these effects are very limited owing to an extremely small number of examined events.

On December 18, 2018 a space body entered the Earth's atmosphere and exploded at the altitude of ~ 25 km. Later it became known as Kamchatka meteor or Bering Sea Bolide. The explosion, which was estimated to be ~ 10 times more powerful than the atomic bomb explosion in Hiroshima ($\sim 21~\rm kT$), was not seen by anyone. It was discovered only as a result of post-processing of photos taken by NASA satellites.

The purpose of our work is to examine possible electromagnetic effects, which may be caused by interaction of a cosmic body with the magnetosphere-ionosphere-atmosphere system. In search for possible effects, we consider not only the final explosion and acoustic blast, but disturbances during passage of the bolide through the inner magnetosphere and ionosphere.

The Bering Sea Event

According to NASA fireball database [https://cneos.jpl.nasa.gov/fireballs] the explosion occurred on December 18, 2018 at 23:48:20 UT in the point with geographic coordinates 56.9° N and 172.4° E. The radiative energy was estimated to be 31 kT and the total energy was about 173 kT. According to this estimate, the Bering Sea Meteoroid (BSM) explosion is the third in power after the Tunguska event ($\sim 20~000~kT$) and the Chelyabinsk meteoroid ($\sim 440~kT$) [Popova et al., 2013], but much more powerful than the Vitim meteoroid ($\sim 2~kT$). The BSM was estimated to have a core diameter $\sim 10~meters$ and weighed $\sim 1~360~tons$.

NASA experts estimated the speed of BSM en-

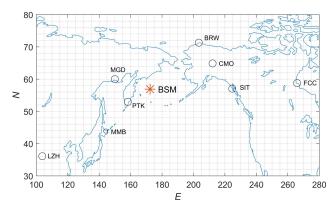


Figure 1. The location of the BSM explosion and the most important geomagnetic stations in the area of the explosion.

try of ~ 32 km/s, with the speed components Vx=6.3, Vy=-3 and Vz=-31.2 km/s. Thus, the BSM entered the atmosphere almost vertically. Therefore, the BSM entered the inner magnetosphere/plasmasphere (~ 3 –4 R_E) about 10–13 minutes before the explosion.

In addition to optical data, acoustic observations at the international monitoring system (IMS) are available. IMS operates in compliance with the comprehensive nuclear-test-ban treaty (CTBT) and includes ~ 60 infrasound stations throughout the globe. Infrasound response to the BSM explosion was identified at distances from ~ 1000 to ~ 8000 km at 13 stations located at azimuths from 17° to 349° [Negraru and Johnson, 2019]. The period of registered acoustic signals was in the range from 2.4 to 20.1 seconds. The equivalent energy release for the bolide may be estimated from the relationship $\log Y = 3.34 \log \tau - 2.58$, where Y is the yield (in kT) and τ is the dominant signal period (in sec) [Revelle, 1997]. As a result, the bolide energy was estimated to be between 4 and 35 kT. This estimate is significantly lower than those obtained from the optical data. The same conclusion was reached by *Gordeev et al.* [2019] based on Russian infrasound measurements in Kamchatka. According to their data, the TNT equivalent of the bolide is estimated to be even lower, about 1.3–8.2 kT. Hence, the explosion power of 173 kT derived from the NASA optical observations seems to be overestimated, since in this case the period of infrasound oscillations should have been at least 30 s [Negrary and Johnson, 2019].

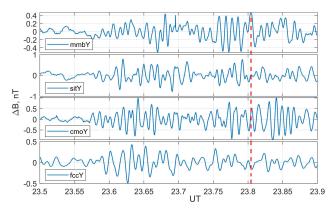


Figure 2. Variations of Y components of the magnetic field according to data of the Memambetsu (MMB), Sitka (SIT), College (CMO) and Fort Churchill (FCC) observatories on December 18, 2018. The red dashed line shows the moment of the explosion.

Geomagnetic Observations

To detect and investigate possible disturbances of the geomagnetic field, we have selected fluxgate magnetometer data from available INTERMAG-NET stations with 1-sec cadence that are located in the vicinity of the explosion site in the Northern hemisphere and its conjugate point in the Southern hemisphere. Besides fluxgate magnetometers from the INTERMAGNET array, we examine the data from more sensitive search-coil magnetometers at Magadan (MGD) and Paratunka (PET) stations with 64 Hz sampling frequency deployed at subauroral latitudes within the framework of the PWING project [Shiokawa et al., 2017]. Figure 1 shows the location of the explosion and position of the most important stations, whereas Table 1 provides information on coordinates of the magnetic observatories considered and their distances R from the explosion site. All stations are in the Sun-illuminated or twilight zone during the event [https://www. timeanddate.com/worldclock/sunearth.html].

Figure 2 shows variations of the magnetic field at several magnetic stations in the region of BSM explosion: Memambetsu (MMB), Sitka (SIT), College (CMO), and Fort Churchill (FCC). The geomagnetic effect of the explosion itself and any aftereffects were not detected. About 12 minutes before the explosion (since 23:36 UT) periodic oscillations emerge that last for ~ 15 min. The largest signal peak-to-peak amplitude ~ 1 nT is observed in

Table 1. Magnetic stations whose data is used in the work

Name	Lat. N (°)	Long. E (°)	Distance to ex- plosion R (km)
Alice Springs	-23.77	133.88	9640
Barrow	71.32	203.38	2160
Cocos-Keeling Islands	-12.188	96.834	10275
College	64.87	212.14	2290
Canberra	-35.32	149.36	10450
Eyrewell	-43.474	172.39	10130
Fort Churchill	58.759	265.91	5096
Macquarie Island	-54.5	159.9	12425
Magadan	60	150	1300
Memambetsu	43.91	144.19	2440
Paratunka	52.97	158.25	1000
Sitka	57.06	224.67	3100
Lanzhou	36.1	103.84	5450
	Alice Springs Barrow Cocos-Keeling Islands College Canberra Eyrewell Fort Churchill Macquarie Island Magadan Memambetsu Paratunka Sitka	Name N (°) Alice Springs -23.77 Barrow 71.32 Cocos-Keeling Islands -12.188 College 64.87 Canberra -35.32 Eyrewell -43.474 Fort Churchill 58.759 Macquarie Island -54.5 Magadan 60 Memambetsu 43.91 Paratunka 52.97 Sitka 57.06	Name N (°) E (°) Alice Springs -23.77 133.88 Barrow 71.32 203.38 Cocos-Keeling Islands -12.188 96.834 College 64.87 212.14 Canberra -35.32 149.36 Eyrewell -43.474 172.39 Fort Churchill 58.759 265.91 Macquarie Island -54.5 159.9 Magadan 60 150 Memambetsu 43.91 144.19 Paratunka 52.97 158.25 Sitka 57.06 224.67

the Y component (E-W). The observed signals are similar to Pc3 geomagnetic pulsations. Typical Pc3 pulsations is a dayside/morning phenomenon, they last for \sim 2–4 hours, and are observed predominantly in the X (N-S) component. In contrast to them, oscillations detected before the BSM explosions are short-lived, and are more evident in Y component.

Figure 3 shows the result of time-frequency analysis (sonogram) of the signal Y component recorded at selected magnetic stations. These spectrograms confirm the appearance of magnetic field pulsations before the explosion. The dominant oscillation frequency is $\sim 25\text{--}35$ mHz (period is ~ 0.5 min). This period corresponds to typical resonant eigenfrequencies of the magnetospheric field lines at the latitudes under study.

More sensitive search-coil magnetometers at MGD and PET confirm the occurrence of weak oscillations with peak-to-peak amplitude $\sim 0.005 \text{ nT/s}$ in the Pc3 band before the explosion (Figure 4).

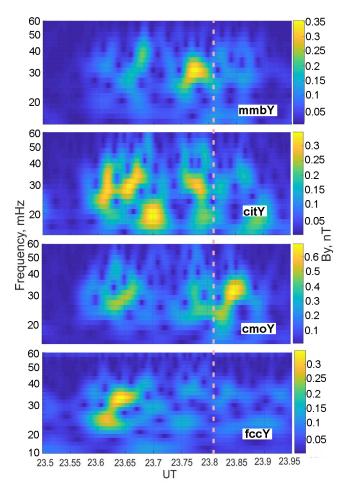


Figure 3. The result of time-frequency analysis of the signal for Y components of the magnetic field for MMB, SIT, CMO and FCC observatories. The red dashed line shows the moment of the explosion.

A clear wave packet appeared in Y component ~ 3 min before the explosion and disappeared just after it

At stations located far from the explosion site (about $5 \cdot 10^3$ km and father), e.g. Lanzhou (LZH) to the West and Ottawa (OTT) to the East, oscillations are practically undetectable. Taking into account the fact that according to NASA database [https://cneos.jpl.nasa.gov/fireballs], the horizontal component of the BSM speed was 6.3 km/s, during the registration of geomagnetic variations the horizontal projection of its trajectory did not exceed $5 \cdot 10^3$ km. Thus, the most intense quasiharmonic 30-s geomagnetic fluctuations have been observed at stations located inside the magnetic tube through which the meteoroid entered.

If one suggests that during the interaction of

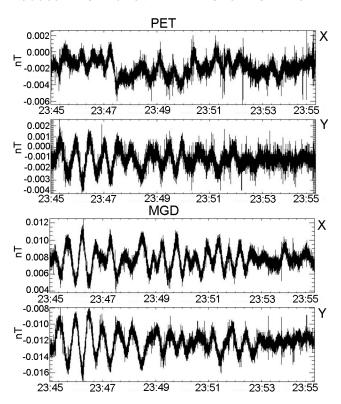


Figure 4. Variations of X and Y components of the magnetic field according to data of the Paratunka (PET) and Magadan (MGD) observatories on December 18, 2018.

the BSM with the magnetospheric plasma the field line oscillations are excited, then the occurrence of similar oscillation train is to be expected in magnetically conjugated point in the Southern hemisphere. To test this assumption, records of geomagnetic variations at Alice Springs (ASP), Canberra (CNB), Cocos-Keeling Islands (CKI) and Macquarie Island (MCQ) in Australia, as well as Eyrewell (EYR) in New Zealand, have been examined. The location of this stations is shown in Figure 5. Spectrograms of magnetic field oscillations from these stations are shown in Figure 6. It can be seen that small-amplitude oscillations similar to those found in Northern hemisphere can be seen only at two stations, CNB and EYR, conjugate to the explosion flux tube. Thus, the effect is localized also in the Southern hemisphere in the region magnetically conjugated with the meteoroid explosion site. Figure 7 shows the result of comparing of magnetic oscillations in the conjugate points of the two hemispheres.

Discussion

A magnetic effect of a meteor can be produced by the ionization along its trail and ionospheric electric current modified by this ionization. The sudden creation of extra ionization and conductivity is equivalent to the superposition of an additional current system whose magnetic field may be observable at the ground. The radar detectable ionization by meteors is usually produced between 80 and 110 km. The theoretical order-of-magnitude estimation of a meteor magnetic effect can be made with a simple model as a vertical uniform cylinder with increased conductivity Σ_1 immersed into a uniform ionosphere with conductivity Σ_0 bounded by horizontal planes at heights h and h + d, and background ionospheric current density j_o [Chapman and Ashour, 1965. The total excess current J flowing through the cylinder with diameter a produces the magnetic response on the ground beneath the meteor trail which can be estimated as follows:

$$J = 2j_0 a d \frac{\Sigma_1 - \Sigma_0}{\Sigma_1 + \Sigma_0} \qquad \Delta B = \frac{\mu_0 J a}{2h(h+d)} .$$

For parameters h=100 km, d=30 km, $j_o=10^{-6}$ A/m², $\Sigma_1/\Sigma_0=3$, a=30 m, this estimate gives the expected magnetic effect $\Delta B\sim 2$ nT. However, no geomagnetic response has been detected when the BSM was in the ionospheric E-layer. Probably, the lapse time when the BSM crossed the conductive E-layer, <1 s, was too short to excite prolonged oscillations.

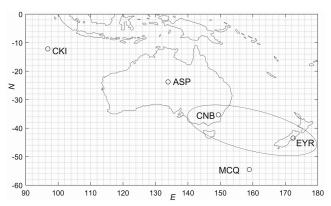


Figure 5. The position of the Alice Springs (ASP), Canberra (CNB), Cocos Islands (CKI) and Macquarie island (MCQ) observatories in Australia and Airwell (EYR) in New Zealand in the area magnetically coupled with BSM explosion of the southern hemisphere.

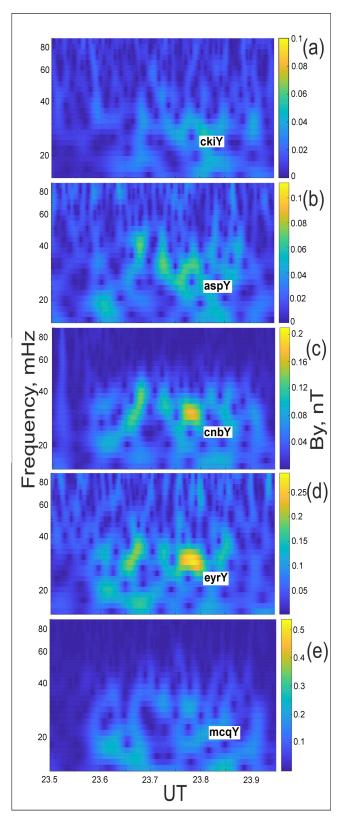


Figure 6. Spectrograms of magnetic field oscillations for CKI (a), ASP (b), CNB (c), EYR (d) and MCQ (e) stations. The red dashed line shows the moment of the explosion.

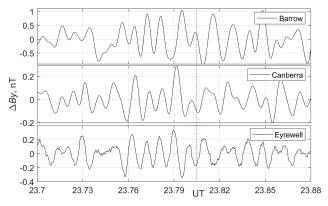


Figure 7. The result of comparing of magnetic oscillations for BRW, CNB and EYR stations in the magnetically conjugate regions of the two hemispheres.

Another electromagnetic effect which could cause the excitation of geomagnetic pulsations is associated with the movement of a conducting body with velocity V_0 through a magnetized plasma is the formation of the Alfven wings [Dobrowolny and Veltri, 1986]. A front of field-aligned currents moves away from the body at an angle with geomagnetic field $\theta = \arctan(V_0/V_A)$, where V_A is the Alfven velocity. The current flowing along an Alfven wing is $J = 4[V_0 \times B]a\Sigma_A$, where a is the scale of a body. The factor $\Sigma_A = (\mu V_A)^{-1}$ is the Alfven wave conductance. This factor is larger inside the plasmasphere, so the Alfven wing generation is more efficient in this domain.

Chernogor [2018] interpreted the geomagnetic field perturbations observed before the Chelyabinsk Meteoroid impact as a result of magnetic field expulsion from the meteoroid trail. To achieve an agreement with ground observations, he had to assume that the magnetospheric magnetic field was completely expelled from an extended cavity with length $\sim 1.5R_E$ and diameter $\sim 2R_E$. A diamagnetic cavity of this size could only be formed if all the kinetic energy of the bolide was used to transfer the meteoroid substance to the plasma state. Such a situation is obviously impossible for BSM at a distance of several thousand kilometers from Earth. However, due to partial ablation of the meteorite, the diamagnetic effect could cause a local disturbance of the geomagnetic field and its propagation in the magnetic force tube.

We must take into account that the magnetosphere is often in a metastable state, when even

a relatively weak external trigger can stimulate an internal instability and wave generation. As an example, one may recall the triggering by solar wind pressure pulses of magnetospheric long-lasting Pc1 emissions or global Pc5 waves [*Pilipenko*, 2013]. In line with this, a possibility of bolide flight to be an external trigger of resonant field line oscillations cannot be excluded. Although the presented evidences indicate that the recorded burst of oscillations in the frequency range around the eigenfrequency of magnetospheric field lines at a latitude under consideration are indeed triggered by a bolide transition through the inner magnetosphere, a possibility of accidental coincidence cannot be absolutely excluded. Nonetheless, keeping in mind a rare occurrence of such events, all associated effects must be examined.

Conclusion

The purpose of this study is to consider electromagnetic phenomena that could be caused by the fall of a large meteoroid over the Bering sea. It is surprising that despite a rather significant scale of the event, it did not attract much attention of the scientific community. There is a certain discrepancy in the explosion energy of this meteoroid, which is estimated from the optical data as 173 kT, and rather weak acoustic and electromagnetic effects that it caused. Unexpectedly, we found that the magnetic signal is recorded before the moment of explosion, so the signal may be generated in the inner magnetosphere during the passage of the meteoroid. The found excitation of magnetic pulsations ~ 10 minutes before the meteoroid explosion and an absence of electromagnetic signals directly associated with the explosion is unusual for such events. Our main purpose is to draw attention to this problem. Further research on these and other manifestations of the entry of a large space body into the Earth's ionosphere and magnetosphere are necessary. More complete set of observational effects will help to achieve an adequate understanding of the physical phenomena associated with meteorites, to improve methods for detecting and evaluating their characteristics, and to identify approaches to the problem of meteoroid danger prevention.

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